

**Title:** Analysis of the pressure at a vertical barrier due to extreme wave run-up over variable bathymetry

**Author(s):** J. Brennan<sup>1</sup>, C. Clancy<sup>1</sup>, J. Harrington<sup>2</sup>, R. Cox<sup>2</sup>, F. Dias<sup>1</sup>

<sup>1</sup>School of Mathematics and Statistics, University College Dublin, Belfield, Dublin 4, Ireland <sup>2</sup>Department of Geosciences, Williams College, Williamstown, MA 01267, USA

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#### Letter

## Analysis of the pressure at a vertical barrier due to extreme wave run-up over variable bathymetry



J. Brennan<sup>a</sup>, C. Clancy<sup>a</sup>, J. Harrington<sup>b</sup>, R. Cox<sup>b</sup>, F. Dias<sup>a,\*</sup>

- <sup>a</sup> School of Mathematics and Statistics, University College Dublin, Belfield, Dublin 4, Ireland
- <sup>b</sup> Department of Geosciences, Williams College, Williamstown, MA 01267, USA

#### HIGHLIGHTS

• Ocean waves may be amplified as they approach coasts across steep bathymetry, leading to extreme run-up heights.

#### ARTICLE INFO

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#### ABSTRACT

The pressure load at a vertical barrier caused by extreme wave run-up is analysed numerically, using the conformal mapping method to solve the two-dimensional free surface Euler equations in a pseudo-spectral model. Previously this problem has been examined in the case of a flat-bottomed geometry. Here, the model is extended to consider a varying bathymetry. Numerical experiments show that an increasing step-like bottom profile may enhance the extreme run-up of long waves but result in a reduced pressure load.

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The existence of extreme waves is an important problem to various fields and has deservedly received considerable attention, in particular on the open ocean where there are many documented cases which have caused loss of life as well as severe damage to infrastructure (see, for example, O'Brien et al. [1] and references mentioned within). As such, in the engineering of both coastal and marine infrastructure, one is provided an interesting challenge when providing durable and safe design. In fact, the engineer often relies on the design wave for estimating various quantities (such as wave height and pressure load) which the structure may experience. Needless to say, this is a difficult task which is made worse by the occurrence of large waves.

A particular case of extreme wave action is the enhanced run-up amplification of long ocean waves impinging on a vertical barrier. The physical mechanisms involved are discussed in detail in the review of Peregrine [2]. The problem has also been studied using numerical simulations by a number of authors. By solving the Euler equations with a conformal mapping model, Carbone et al. [3] and Viotti et al. [4] explored the nonlinear evolution of the wave run-up phenomenon. They showed that a simple, nearly monochromatic

wave group can produce a run-up amplitude in excess of six times the far-field amplitude. Viotti et al. [5] extended the model to account for non-flat bathymetry, while Brennan et al. [6] examined the associated pressure fluctuations on the barrier. Akrish et al. [7,8] reproduced the run-up using a higher order spectral (HOS) method, and also looked at the resulting force on the wall.

In this present work, we wish to extend the analysis of the pressure load on the barrier to the case of a varying bathymetry. We begin by describing the mathematical formulation of the conformal mapping method for the Euler equations along with details of the numerical model used. Under the assumption of inviscid, irrotational fluid flow, the dynamics of the associated free surface wave system are described by the Euler equations:

$$\nabla^2 \phi = 0, \text{ for } b(x) < y < \zeta(x, t), \ 0 < x < l, \tag{1}$$

$$\zeta_t = -\phi_x \zeta_x + \phi_y, \text{ at } y = \zeta(x, t), \tag{2}$$

$$\phi_t = -\frac{1}{2} \left( \phi_x^2 + \phi_y^2 \right) - g\zeta, \text{ at } y = \zeta(x, t),$$
 (3)

$$\nabla \phi \cdot \mathbf{n} = 0, \text{ at } y = b(x), \tag{4}$$

where  $\phi(x, y, t)$  is the velocity potential,  $\zeta(x, t)$  is the free surface elevation with respect to the unperturbed condition, g is the vertical acceleration due to gravity, and b(x) is the bathymetry profile.

<sup>\*</sup> Corresponding author. E-mail address: frederic.dias@ucd.ie (F. Dias).

Without loss of generality, we shall operate in non-dimensional units in which g is equal to one, and the fluid density is understood to be unitary. The conformal mapping method for solving the Euler equations, originally presented by Dyachenko et al. [9], and later analysed by Choi and Camassa [10], yields time evolution equations for the free surface, and the velocity potential evaluated at the free surface, of the gravity wave system. Viotti et al. [5] applied conformal mapping to a system containing a varying bottom. Presented here within is a brief overview of the formulation, highlighting the features important to the discussion. The technique involves finding an appropriate analytical map  $Z = X(\xi, \eta, t) +$  $iY(\xi, \eta, t)$  between the physical domain (x, y) and the conformal domain  $(\xi, \eta)$ . Here, we make the distinction that capitalisation of variables implies that they are functions of conformal variables. The physical domain is naturally defined by the fluid in question, while the conformal domain is described by an arbitrary strip of uniform thickness H > 0.

Boundary conditions for the harmonic function  $Y(\xi, \eta, t)$  are provided by the vertical displacement of the free surface and the given bathymetry profile:

$$Y = \tilde{Y}(\xi, t), \text{ at } \eta = 0,$$
  

$$Y = B(\xi, t), \text{ at } \eta = -H,$$
(5)

which, assuming periodicity of the problem in  $\xi$  with period  $L = 2\pi/k$ , we expand as

$$\tilde{Y}(\xi, t) = \sum_{n \neq 0} \hat{Y}_n e^{ink\xi}, 
B(\xi, t) = \sum_{n \neq 0} \hat{B}_n e^{ink\xi},$$
(6)

where  $\hat{Y}_n$  and  $\hat{B}_n$  are the Fourier coefficients of the free surface and bathymetry, respectively. Note that we are implicitly assuming a periodic bathymetric profile in this case. Furthermore, we have introduced the notation that tildes represent variables restricted to the free surface. The most general solution to these conditions is

$$Y(\xi, \eta, t) = \hat{Y}_0 + \left(\frac{\hat{Y}_0 - \hat{B}_0}{H}\right) \eta + \sum_{n \neq 0} \hat{Y}_n \frac{\sinh\left[nk(\eta + H)\right]}{\sinh(nkH)} e^{ink\xi}$$
$$- \sum_{n \neq 0} \hat{B}_n \frac{\sinh(nk\eta)}{\sinh(nkH)} e^{ink\xi}. \tag{7}$$

Given the harmonic nature of the problem, we take advantage of the Cauchy–Riemann conditions to derive the corresponding expression for  $X(\xi, \eta, t)$ ,

$$X(\xi, \eta, t) = x_0(t) + \left(\frac{\hat{Y}_0 - \hat{B}_0}{H}\right) \xi - i \sum_{n \neq 0} \hat{Y}_n \frac{\cosh[nk(\eta + H)]}{\sinh(nkH)} e^{ink\xi} + i \sum_{n \neq 0} \hat{B}_n \frac{\cosh(nk\eta)}{\sinh(nkH)} e^{ink\xi},$$
(8)

where  $x_0(t)$  is an integration constant. Here, we note that if the expression  $H = \hat{Y}_0 - \hat{B}_0$  holds, then the first Fourier coefficient of  $Y_\eta$  (=  $X_\xi$ ) is equal to one. As such we assume this, thereby setting the spatial period in conformal space and physical space equal (i.e., L = I). This also implies that H is time dependent, which is a natural consequence of the conformal mapping. However, the physical bathymetry profile b(x) has no dependence on time.

Similarly, we can find expansions for the velocity potential  $\Phi$  and the stream function  $\Psi$ :

$$\Phi(\xi, \eta, t) = \mathcal{U}\xi + \sum_{n \neq 0} \hat{\Phi}_n \frac{\cosh[nk(\eta + H)]}{\cosh(nkH)} e^{ink\xi}, \tag{9}$$

$$\Psi(\xi, \eta, t) = \mathcal{U}(\eta + H) + i \sum_{n \neq 0} \hat{\Phi}_n \frac{\sinh[nk(\eta + H)]}{\cosh(nkH)} e^{ink\xi}, \quad (10)$$

where  $\hat{\Phi}_n$  are the Fourier coefficients of the velocity potential on the free surface, and  $\mathcal{U}=m[\tilde{\Psi}]/H$  is the mean horizontal velocity of the flow (the operator m refers to integral mean over one period), included so that the periodic nature of the potential may be isolated:

$$\tilde{\Phi}(\xi,t) - \mathcal{U}\xi = \sum_{n \neq 0} \hat{\Phi}_n e^{ink\xi}.$$
 (11)

The Hilbert-like transform operators provide a simple means of transforming between the various harmonic conjugate pairs in conformal space. On the free surface, the following relations hold

$$\tilde{X}_{\xi} = 1 - \mathcal{T}_{\chi}[\tilde{Y}_{\xi}, B_{\xi}], \quad \tilde{\Psi}_{\xi} = -\mathcal{T}_{\psi}[\tilde{\Phi}_{\xi} - \mathcal{U}], \tag{12}$$

with  $\mathcal{T}_x$  and  $\mathcal{T}_{\psi}$  operating as follows:

$$\mathcal{T}_{x}[f,g] = \frac{1}{H} \int f(\xi') \coth[\pi(\xi'-\xi)/2H] d\xi' + \frac{1}{H} \int (g(\xi')-H) \tanh[\pi(\xi'-\xi)/2H] d\xi',$$
(13)

$$\mathcal{T}_{\psi}[f] = -\frac{1}{H} \int f(\xi') \operatorname{cosech}[\pi(\xi' - \xi)/2H] d\xi'. \tag{14}$$

Note that f refers to the principal-value integral over the real axis. We are now in position to express the system's governing equations in conformal space (see Refs. [5] and [10] for a thorough derivation). The time evolution of the free surface is given by

$$\tilde{Y}_{t} = -\tilde{X}_{\xi} \left( \frac{\tilde{\Psi}_{\xi}}{J} \right) + \tilde{Y}_{\xi} \mathcal{T}_{x} \left[ \frac{\tilde{\Psi}_{\xi}}{J}, H_{t} \right] + \tilde{Y}_{\xi} q(t), \tag{15}$$

where  $J = \tilde{X}_{\xi}^2 + \tilde{Y}_{\xi}^2$  is the Jacobian on the free surface, and q(t) is related to the constant  $x_0(t)$  present in Eq. (8). Before defining q(t), it is important to note that  $H_t$  does not depend on  $\xi$ , and thus will produce only a constant value under the action of  $\mathcal{T}_x$  transform. As such, we may absorb the effect of  $H_t$  in q(t), and set it to zero within the transform. We set q(t) as

$$q(t) = m \left\{ X_{\xi} \mathcal{T}_{x} \left[ \frac{\psi_{\xi}}{J}, 0 \right] + Y_{\xi} \left( \frac{\psi_{\xi}}{J} \right) \right\}, \tag{16}$$

a definition which implies that  $x_0$  is a constant in time. Finally, the evolution equations of the system are

$$\tilde{Y}_{t} = -\tilde{X}_{\xi} \left( \frac{\tilde{\Psi}_{\xi}}{J} \right) + \tilde{Y}_{\xi} \mathcal{T}_{x} \left[ \frac{\tilde{\Psi}_{\xi}}{J}, 0 \right] + \tilde{Y}_{\xi} q(t), \tag{17}$$

$$\tilde{\Phi}_t = -\frac{1}{J} \left\{ \frac{1}{2} \tilde{\Phi}_{\xi}^2 - \frac{1}{2} \tilde{\Psi}_{\xi}^2 - J \tilde{\Phi}_{\xi} \mathcal{T}_{\chi} \left[ \frac{\tilde{\Psi}_{\xi}}{J}, 0 \right] \right\} - g \tilde{Y} + C(t). \quad (18)$$

We note that C(t) is an arbitrary function of time which may be absorbed into  $\tilde{\Phi}_t$ .

The Bernoulli equation may be exploited in order to calculate the total pressure associated with the system in question. Note that this was done previously in Ref. [6] for the special case of flat bathymetry. Here, we extend to the varying sea bed. The Bernoulli equation is given by

$$P = -\left[\phi_t + \frac{1}{2}\left(\phi_x^2 + \phi_y^2\right) + gy\right] + D,\tag{19}$$

where D is an arbitrary constant. The required derivatives of  $\phi$  necessary for the computation of Eq. (19) can be determined by evaluating  $\Phi_{\xi}$ ,  $\Phi_{\eta}$ , and  $\Phi_{t}$  via the chain rule, and rearranging the resulting equations:

$$\phi_{x} = \frac{\Phi_{\xi} - \Phi_{\eta} \frac{Y_{\xi}}{Y_{\eta}}}{X_{\xi} - X_{\eta} \frac{Y_{\xi}}{Y_{\eta}}}, \quad \phi_{y} = \frac{\Phi_{\eta} - \phi_{x} X_{\eta}}{Y_{\eta}}.$$
 (20)

$$\phi_t = \Phi_t - \phi_x X_t - \phi_y Y_t. \tag{21}$$

While  $\phi_x$  and  $\phi_y$  are easily found via the Fourier expansion equations,  $\phi_t$ , however, requires more thought. First, we note that the conformal mapping procedure conveniently provides the Fourier coefficients of  $Y_t$  and  $\Phi_t$  on the free surface,  $\hat{Y}_{n_t}$  and  $\hat{\Phi}_{n_t}$ . As  $(X_t, Y_t)$  and  $(\Phi_y, \Psi_t)$  are harmonic conjugate pairs, we seek to express  $\hat{X}_t$  in terms of  $\hat{Y}_t$ , which is achieved via the operator  $\mathcal{T}_x$ ;  $\hat{X}_{n_t} = -\frac{\partial}{\partial t}\mathcal{T}_x\left[\hat{Y}_n,\hat{B}_n\right]$ .

This provides the means to construct the time derivative of the analytic map over the whole domain,  $Z_t$ . On the free surface, the map is given by

$$\tilde{Z}_t(\xi,t) = \sum_{n \neq 0} \left( \hat{X}_{n_t} e^{ink\xi} + i\hat{Y}_{n_t} e^{ink\xi} \right), \tag{22}$$

whose analytical extension to the full domain is

$$Z_{t}(\xi, \eta, t) = \sum_{n \neq 0} \left( \hat{X}_{n_{t}} e^{ink(\xi + i\eta)} + i\hat{Y}_{n_{t}} e^{ink(\xi + i\eta)} \right)$$
$$= \sum_{n \neq 0} \hat{Z}_{n_{t}} e^{ink\xi}. \tag{23}$$

 $\hat{Z}_{n_t}$  is given by

$$\begin{split} \hat{Z}_{n_{t}} &= \left\{ -\frac{\partial}{\partial t} \mathcal{T}_{x} \left[ \hat{Y}_{n}, \hat{B}_{n} \right] + i \hat{Y}_{n_{t}} \right\} e^{-nk\eta} \\ &= \left[ -i \hat{Y}_{n_{t}} \coth(nkH) \right. \\ &+ i nkH_{t} \left( \frac{1}{\sinh^{2}(nkH)} \hat{Y}_{n} - \coth(nkH) \operatorname{cosech}(nkH) \hat{B}_{n} \right) \\ &+ \frac{i}{\sinh(nkH)} \hat{B}_{n_{t}} + i \hat{Y}_{n_{t}} \right] e^{-nk\eta} \\ &= i \left[ \frac{-e^{-nkh}}{\sinh(nkH)} \hat{Y}_{n_{t}} + \frac{nkH_{t}}{\sinh^{2}(nkH)} \left( \hat{Y}_{n} - \cosh(nkH) \hat{B}_{n} \right) \right. \\ &+ \frac{1}{\sinh(nkH)} \hat{B}_{n_{t}} \right] e^{-nk\eta}, \end{split}$$

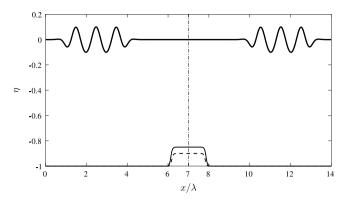
where the rearrangement of terms is performed in order to alleviate truncation errors during numerical computation.  $X_t$  and  $Y_t$  are thus given by

$$X_t = \operatorname{Re}(Z_t), \tag{25}$$

$$Y_t = \operatorname{Im}(Z_t). \tag{26}$$

Finally, as they satisfy the same boundary conditions, we may Fourier expand  $\Phi_t$  in the same manner as  $\Phi$ , and thus all components of Eq. (19) are found.

The system's evolution equations in conformal space, Eqs. (17)–(18) are solved via a pseudo-spectral method, with transforms being handled by the fast Fourier transform. The corresponding discretisation of the system follows standard procedure, and due to the high order of nonlinearity inherent to the problem, dealiasing is performed using a 1/2 rule, and applied sequentially after every nonlinear operation. Time integration is handled via an optimal adaptive ordinary differential equation solver, again to the standard of common software packages (for reference, we use MATLAB).



**Fig. 1.** Initial configuration of numerical simulation. Thick line: Initial free surface. Thinner lines: bathymetry profile, with S=0.10 (dashed line) and S=0.15 (solid line). The dashed vertical line indicates the "barrier", across which mirror symmetry is imposed. Here, the wave packet's initial wave amplitude is  $a_0=0.10$ . Horizontal coordinate x is normalised by initial packet wavelength  $\lambda_0$ .

Mirror symmetry about the centreline of the computational domain is enforced on the initial condition, resulting in two wave packets of equal amplitude and opposite direction, starting from opposite ends of the wave tank, colliding head-on at the centre. In the absence of viscosity and surface tension, this is identical to the presence of a rigid vertical barrier [11–13]. Thus we emulate wave reflection from the vertical barrier by simulating two identical waves colliding. To clarify, the physical domain can be thought of as  $0 \le x \le L$ , the computational domain is  $0 \le x \le 2L$ , and the "barrier" is located at x = L.

The free surface  $\tilde{y}_0(x) = \zeta_0(x)$  and the associated velocity potential  $\tilde{\phi}_0(x)$  have the following initial conditions:

$$\tilde{y}_0(x) = W(x)a_0 \sin[k_0(x - x_0)],$$
(27)

$$\tilde{\phi}_0(x) = -W(x)a_0 \frac{\omega_0}{gk_0} \coth(k_0 h) \cos[k_0(x - x_0)], \qquad (28)$$

where  $\omega_0$  is the initial frequency, satisfying the finite-depth dispersion relation

$$\omega_0^2 = gk_0 \tanh(k_0 h),\tag{29}$$

and W(x) is the envelope function, satisfying

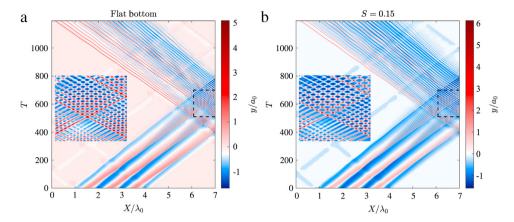
$$W(x) = \frac{1}{2} \left[ \tanh\left(\frac{x - x_0}{\delta}\right) - \tanh\left(\frac{x - x_0 - 3\lambda_0}{\delta}\right) \right]. \tag{30}$$

This particular envelope function produces an initial wave packet containing three full wavelengths, and the initial (dimensionless) wavelength is set as  $\lambda_0=125$ . The parameter determining the location of the initial wave packet  $x_0$  is fixed at one wavelength. Additionally,  $\delta\approx0.3\lambda_0$  is the thickness parameter, set as such to minimise the formation of secondary waves in the early time evolution. These conditions are chosen such as to produce near optimal run-up, see Ref. [4]. The bathymetry profile considered is given by

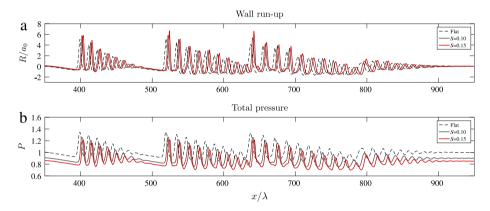
$$b(x) = -h + \frac{S}{2} \left[ \tanh \left( \frac{x - x_1}{\delta_s} \right) - \tanh \left( \frac{x_2 - x}{\delta_s} \right) \right], \tag{31}$$

which produces a smooth ramp followed by flat "step". Here, h is the flat ground level, set equal to one, S is the "step height", and  $\delta_S$  is another thickness parameter which we set equal to L/100, such that in conjunction with the step edge coordinates  $x_1$  and  $x_2$ , produces a "step length" of approximately one wavelength. The domain and set up is depicted in Fig. 1. Simulations are typically performed using  $2^{15}$  (32768) dealiased Fourier modes.

The bottom profile in conformal space,  $B(\xi, t)$ , which depends on both the instantaneous free surface  $\tilde{Y}(\xi, t)$  and the physical



**Fig. 2.** Space time evolution of wave packet, for (a) flat bottom case and (b) step height S=0.15 case. Free surface elevation is normalised by initial dimensionless amplitude. The inset in each panel is a zoomed in snapshot of the regions enclosed by the dashed black lines. Note the differing colour scales used, corresponding to the differing run-up amplitudes found in each case. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)



**Fig. 3.** (a) Run-up measured at wall. (b) Corresponding pressure load on wall. Thick grey lines: Flat bottom case, dashed lines: S = 0.10 case, thin black lines: S = 0.15 case. Note that with a stepped bathymetric profile, there is an increased shoaling effect and run-up amplitude, and conversely, reduced pressure load on wall. The effect is enhanced with increased step height.

bathymetry b(x) needs to be updated at each time step before any computations of the evolution equations can be performed. Further details may be found in Ref. [5]. We utilise a fixed point iteration scheme to update  $B(\xi,t)$  from the previous time step's estimate, with the termination condition being that the residue  $R^{(n+1)} = \|H^{(n+1)} - H^{(n)}\|_{\infty}$  has decreased below a fixed tolerance  $\tau_R$ . The nature of the iteration scheme means that although the physical bathymetry profiles are prescribed to have a step length of one wavelength, this is not exactly guaranteed, and thus we refer to it as an approximate measurement.

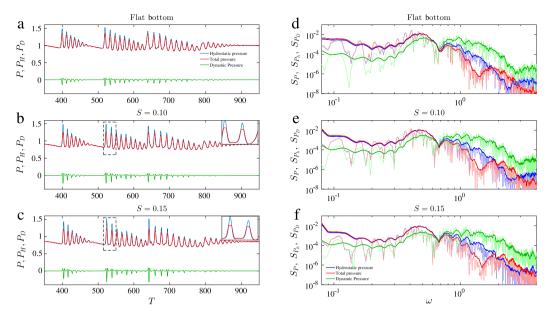
Initially, we examine the effect of varying the step height in the bottom profile. Two step heights are chosen. S=0.10 and S=0.15, and are compared with the flat bottom case. The wave packet has an initial amplitude  $a_0=0.1$ , this value being the interim amplitude examined in Ref. [6]. Time evolution of the wave packet is initially governed by shallow water dynamics, developing sharp wave fronts under the effects of nonlinear steepening. The step bathymetry amplifies this phenomenon. Eventually, dispersive effects regularise the sharp fronts, generating modulated undular bores. Furthermore, enhancement of wave amplification and runup follows as the waves meet the wall. These features are presented in Fig. 2, which also compares the space time evolution of the flat bottom case and S=0.15 case.

The lower panel of Fig. 3 displays the total pressure loads on the wall for each case. The increase in step height is followed by an alleviation of total pressure. This is examined in more detail in Fig. 4, which shows, in addition to total pressure, the hydrostatic and dynamic components, as well as associated power spectra. A

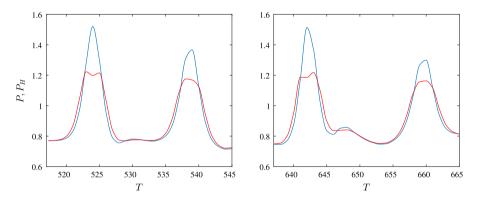
more detailed description of the dynamics of run-up amplification and wave evolution in this scenario can be found in Refs. [3] and [4]. Additionally, similar bore formation can be seen in the recent works of Trillo et al. [14] and Brühl et al. [15], linked with the soliton solutions of the Korteweg–de Vries equation.

Increasing the step height increases the maximum run-up, from  $R/a_0 = 5.23$  for the flat case, to  $R/a_0 = 6.71$  for the maximum step height considered. In each case, the maximum run-up peak is featured in the second run-up packet. There is also a noticeable shoaling effect, which is enhanced with larger step heights. These features can be seen clearly in the upper panel of Fig. 3.

The lower panel of Fig. 3 displays the total pressure loads on the wall for each case. The increase in step height is followed by an alleviation of total pressure. This is examined in more detail in Fig. 4, which shows, in addition to total pressure, the hydrostatic and dynamic components, as well as associated power spectra. The hydrostatic contribution is approximated by  $P_H = -g(y - y)$  $\eta$ ), and dynamical pressure is then  $P-P_H$ . With increasing step height, the reduction in total pressure is mirrored by an increase in the magnitude of the dynamic contribution. With increasing step height, kinks in the pressure peaks begin to appear (these kinks are highlighted in the figure, and are shown in more detail in Fig. 5, a feature which is likely a result of the associated increase in crest height at the wall). These kinks were also observed in Ref. [6], as a result of increasing wave amplitude. The kinks are primarily present in the pressure peaks associated to maximum run-up. Peregrine [2] notes that this is a result of the increasing steepness



**Fig. 4.** (a)–(c) Break down of pressure components, (d)–(f) pressure spectra. Blue lines: hydrostatic pressure, red lines: total pressure, green lines: dynamic pressure. The dynamic contribution to pressure is increased with increasing bathymetric step height. Note that for increased dynamic pressure (reduced total pressure), kinks begin to appear in the pressure peaks associated to maximum run-up, a result of pressure minima forming beneath increasingly steep wave crests. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)



**Fig. 5.** Close up plot of the pressure "kinks" from simulation with step height S = 0.15, from the bottom panel of Fig. 4. Blue line: hydrostatic pressure, red line: total pressure. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

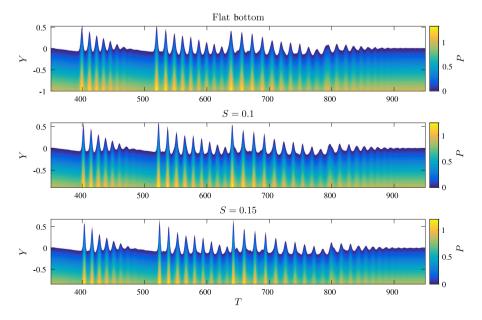
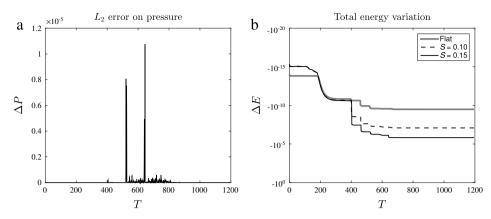
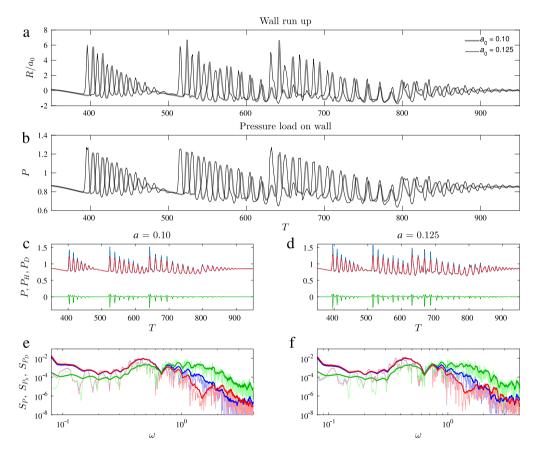


Fig. 6. Total pressure load on wall, including full vertical stratification, for flat and both bathymetric stepped profiles.



**Fig. 7.** (a) Relative  $L_2$  error on total pressure, compared between simulations ran using  $2^{14}$  and  $2^{15}$  Fourier modes, for S=0.15 case. (b) Energy variation for each case. Thick grey line: flat bathymetry, dashed line: S=0.10, black line: S=0.15. Presence of step bathymetry leads to increased energy loss as waves meet vertical barrier.



**Fig. 8.** Examining the effect of varying initial amplitude. Two initial amplitudes,  $a_0 = 0.10$  and  $a_0 = 0.125$ , are chosen. The step height is kept constant at S = 0.15. Colour coding as before. (a) Run-up compared between both cases. (b) Total pressure load on wall, for both cases. The appearance of the pressure "kinks", associated with large run-up peaks, is yet again present. (c)–(f) Comparison between hydrostatic and dynamic pressure components for (c, e)  $a_0 = 0.1$  and (d, f)  $a_0 = 0.125$  cases. Note that by increasing the initial amplitude of the wave packet, similar effects to those induced by increased step height are observed.

of the wave, which creates a pressure minima beneath the wave crest. Examining the power spectrum in Fig. 4, the increasing contribution from dynamic pressure is seen in the high-frequency domain, coinciding with the drop in total pressure compared to the hydrostatic approximation. Interestingly, further into the high-frequency region of the spectrum, the total component briefly overtakes the hydrostatic, at which point the dynamic contribution steadily decreases also, and this effect is diminished with increasing step height. The full pressure fields underneath the wave runup for each case is shown in Fig. 6, where the reduction in total pressure as a result of increasing step height can be seen.

The relative  $L_2$  error on the total pressure load on the wall is shown in the left panel of Fig. 7, computed for the S=0.15 case and compared between two simulations using  $2^{14}$  and  $2^{15}$  dealiased Fourier modes. Feasible convergence is found, with the spikes in the  $L_2$  measurement to order  $O(10^{-5})$  found near the pressure kinks, which may be attributed to the increased nonlinearity of the system as the waves impinge on the wall. Additionally, the total energy variation of the system for each case is shown in the right panel of Fig. 7. Noticeable energy loss is found when the leading peak of each wave group approaches the wall, although reasonable conservation is obtained.

Finally, Fig. 8 displays a comparison between simulation results from two different initial amplitudes. The bathymetry profile is the same in both, using step height S=0.15, and the two initial amplitudes are  $a_0=0.1$  and  $a_0=0.125$ . As can be seen, after the first wave group hits the wall, the run-up for the lower initial amplitude case is actually larger. This is likely related to optimal run-up conditions in the presence of a varying bottom, and is not thoroughly examined in this work. The converse is true for the pressure load, which is larger for the increased amplitude case. Although the run-up is slightly smaller for increased amplitude, the waves themselves are still larger, and as such, the pressure is too. Finally, for the increased amplitude case, the largest run-up is now found in the leading run-up packet. As such, the pressure "kinks" are present again, however given the largest run-up peak is now in the leading packet, the kink is now positioned as such.

In conclusion, we have presented an examination of extreme wave run-up on a vertical wall, in the presence of a varying bathymetry, and in addition, the associated pressure fluctuations are computed. This investigation is facilitated by Euler equation numerical simulations based on the conformal mapping method. The procedure used to identify the pressure field beneath the free surface is also described.

An increasing bottom profile enhances already extreme wave run-up, but decreases the associated pressure load, which is overestimated by the hydrostatic approximation. As such, the dynamic pressure contribution is amplified by the rising sea bed. In terms of the pressure spectrum, this increased dynamic component is seen in the high-frequency region. Increasing the amplitude of the initial wave packet also enhances this phenomenon. The pressure peaks associated with the largest run-up feature the most striking dynamical component, giving rise to the formation of "kinks" in the peaks.

We note that this is a simple investigation, in which we have highlighted the potential effect the varying bottom profile has. A wide class of profiles can be considered, providing scope for future work. The wave parameters considered in this work were chosen to provide near optimal run-up, of which a detailed description can be found in Refs. [3,4]. Optimal conditions in the presence of varying bathymetry have not yet been determined.

As mentioned in Ref. [6], aside from this works relevance to maritime engineering, there is also an interesting implication to microseisms. Longuet–Higgins [16] work on microseisms provided a basis for research on the influence of nonlinear waves on seismic noise, and it is still an active topic today [17,18]. Pressure fluctuations on the sea bed associated with the run-up of long waves contain a considerable high-frequency component in their power spectrum, similar to the typical frequency band associated with microseisms.

A further potential application of this work is in the investigation of boulder-sized deposits along exposed coastlines. The movement of such large masses indicates that the affected coasts experience extremely high-energy wave events (for example, see Cox et al. [19]). Recent research [19,20] focussed on the west coast of Ireland has shown that storm waves, rather than tsunamis, are capable of producing such movements, despite calculations indicating that this should not be possible [21]. This suggests that fully-nonlinear wave models, with the resulting extreme run-up

and pressure loads discussed in this present work, are needed to completely understand these phenomena.

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